

Supplement of Atmos. Meas. Tech., 13, 661–684, 2020
<https://doi.org/10.5194/amt-13-661-2020-supplement>
© Author(s) 2020. This work is distributed under
the Creative Commons Attribution 4.0 License.



Supplement of

Comparison of aircraft measurements during GoAmazon2014/5 and ACRIDICON-CHUVA

Fan Mei et al.

Correspondence to: Fan Mei (fan.mei@pnnl.gov)

The copyright of individual parts of the supplement might differ from the CC BY 4.0 License.

1 Supplemental material:

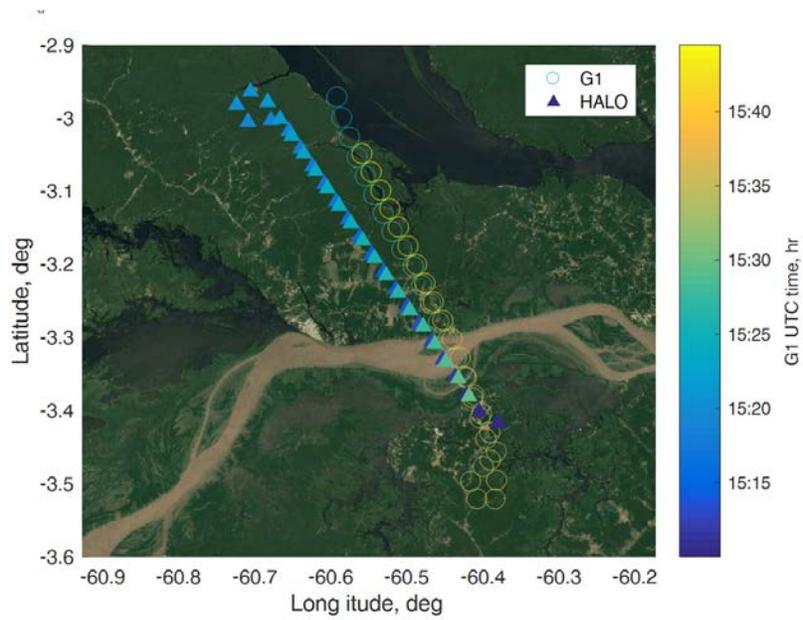
2 **1. Additional observations**

3 Both aircraft cannot appear at the same location at the same time due to safety concerns.
4 Thus, the approval of a formation (inter-comparison) flight was acquired six months before the
5 campaign through DOE Pacific Northwest Site Office (PNSO) and the Office of Aviation
6 Management (OAM). Essential risk mitigation was also discussed and approved by the Pacific
7 Northwest National Laboratory Aviation Risk Management Committee (PNNL ARMC). During
8 the IOP, both aircraft crew and scientists teams set up a meeting to discuss the potential flight plan.
9 After the flight plan was formed, both pilots briefed the plan to the Brazilian Air Force (BAF) and
10 Airport Traffic Control (ATC). The clear-sky flight would be under Visual Flight Rules (VFR),
11 which means good weather and no cloud, and pilots communicate with each other using an air-to-
12 air frequency. For coordinated flights in cloudy conditions, the G1 and the HALO were both on
13 Instrument Flight Rules (IFR) flight plan.

14 The coordinated flight on October 1, 2014, was initially designed to be a coordinated flight
15 under a cloudy condition, which means the G1 and the HALO flew the same flight leg with at least
16 300 m altitude offset and at least 5 minutes apart. However, the coordinated two flight legs (~900
17 m and ~1200 m) are all below the cloud. Thus, the comparison focus on the correlation between
18 two aircraft measurements, not vertical profiling.

19

20

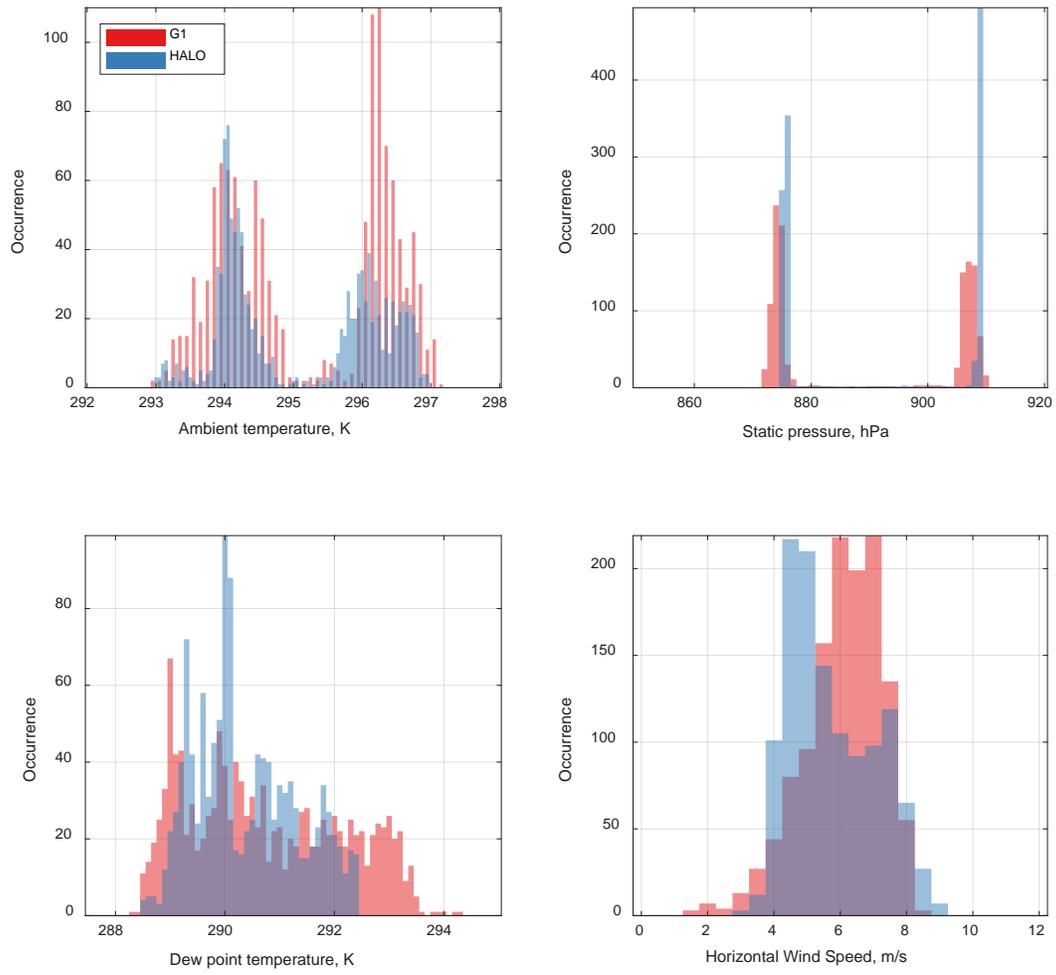


21

22 Figure S1. Time colored flight track of the G1 (circle) and the HALO (triangle) on October 1,
 23 2014, during a cloudless coordinated flight (This figure was created using Mapping Toolbox™

24

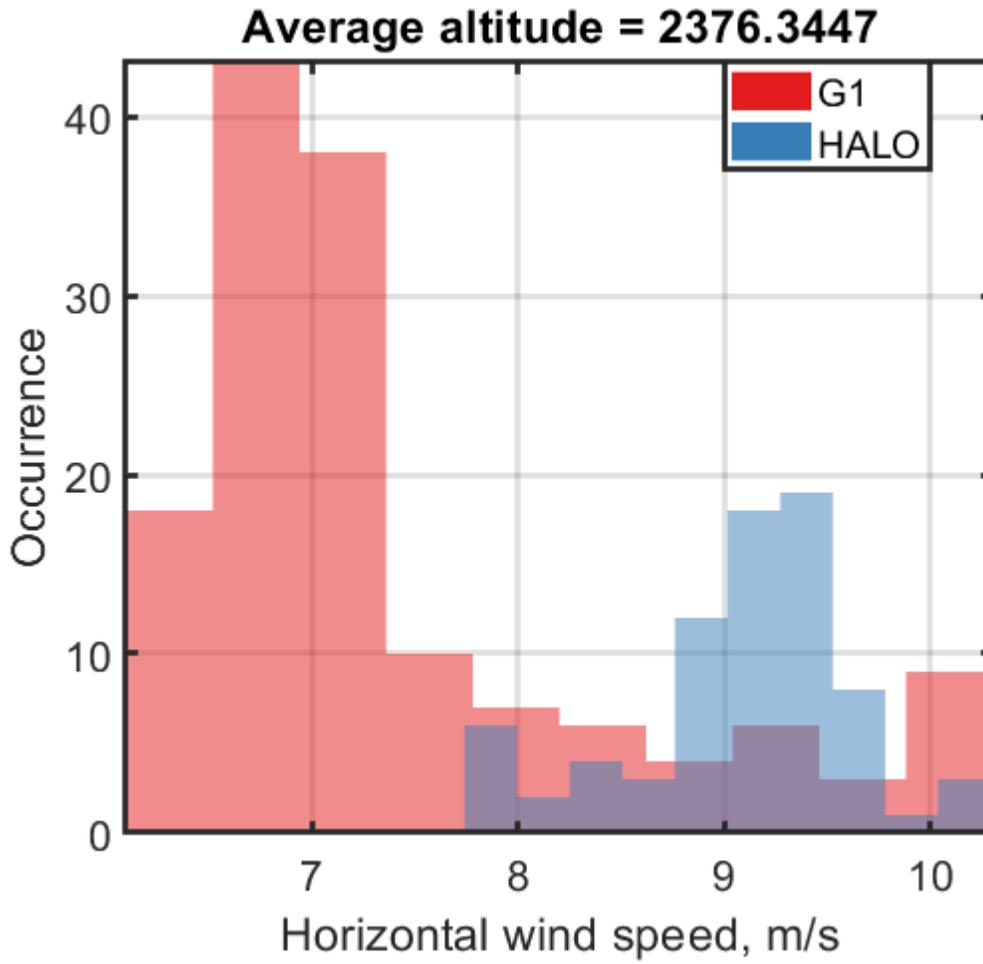
© COPYRIGHT 1997–2019 by The MathWorks, Inc).



25

26

Figure S2, Atmospheric parameters observed by the G1 and the HALO on October 1, 2014.

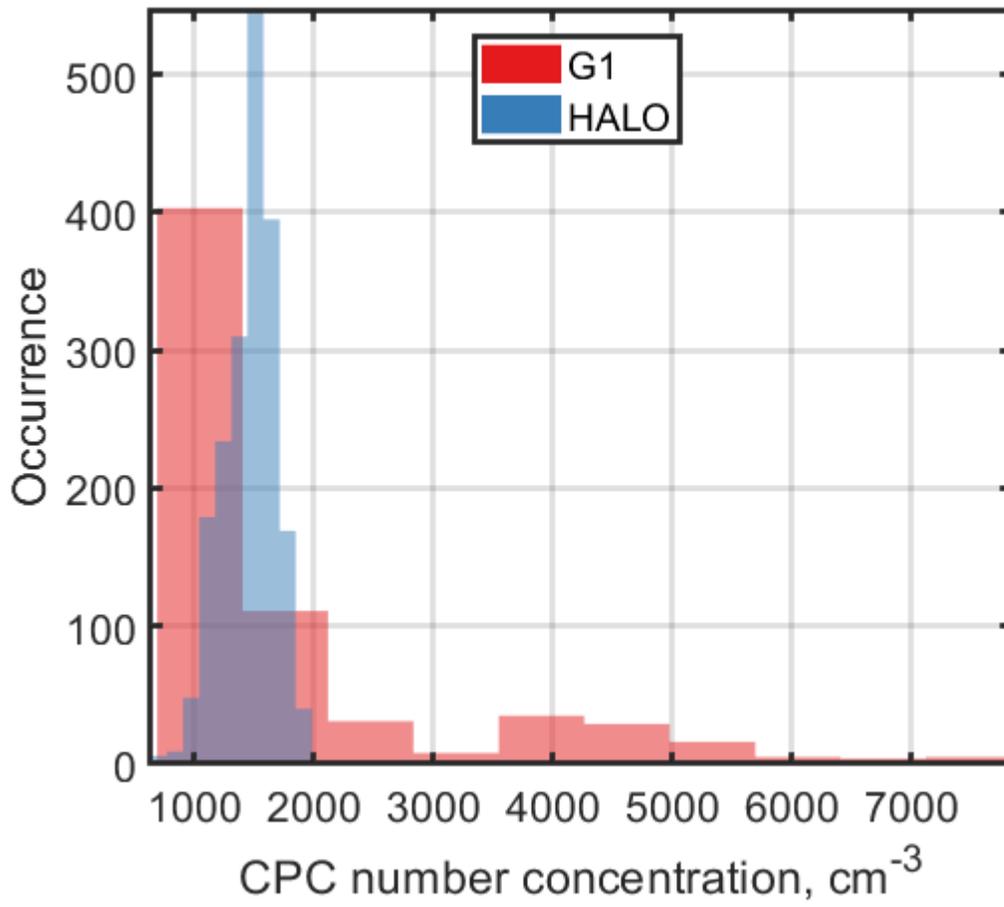


27

28

Figure S3. Horizontal wind speed between 2000-3000 m altitude on September 21, 2014.

29

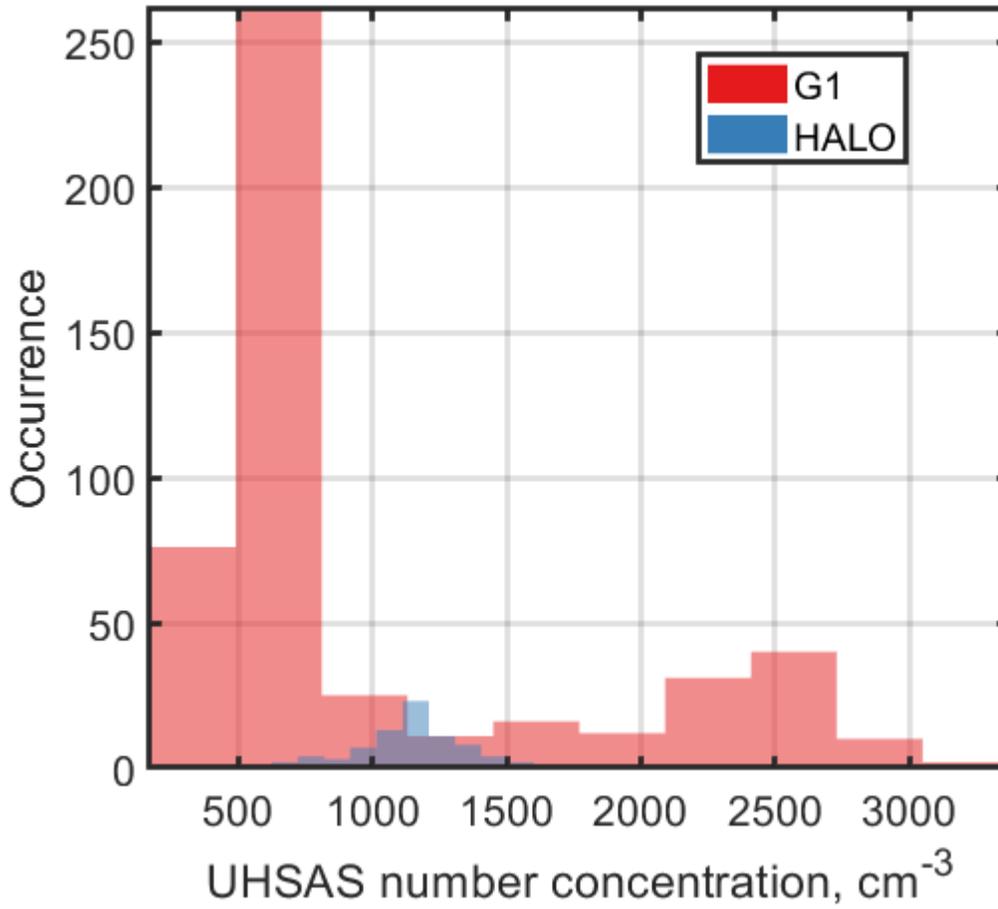


(a)

30

31

32



(b)

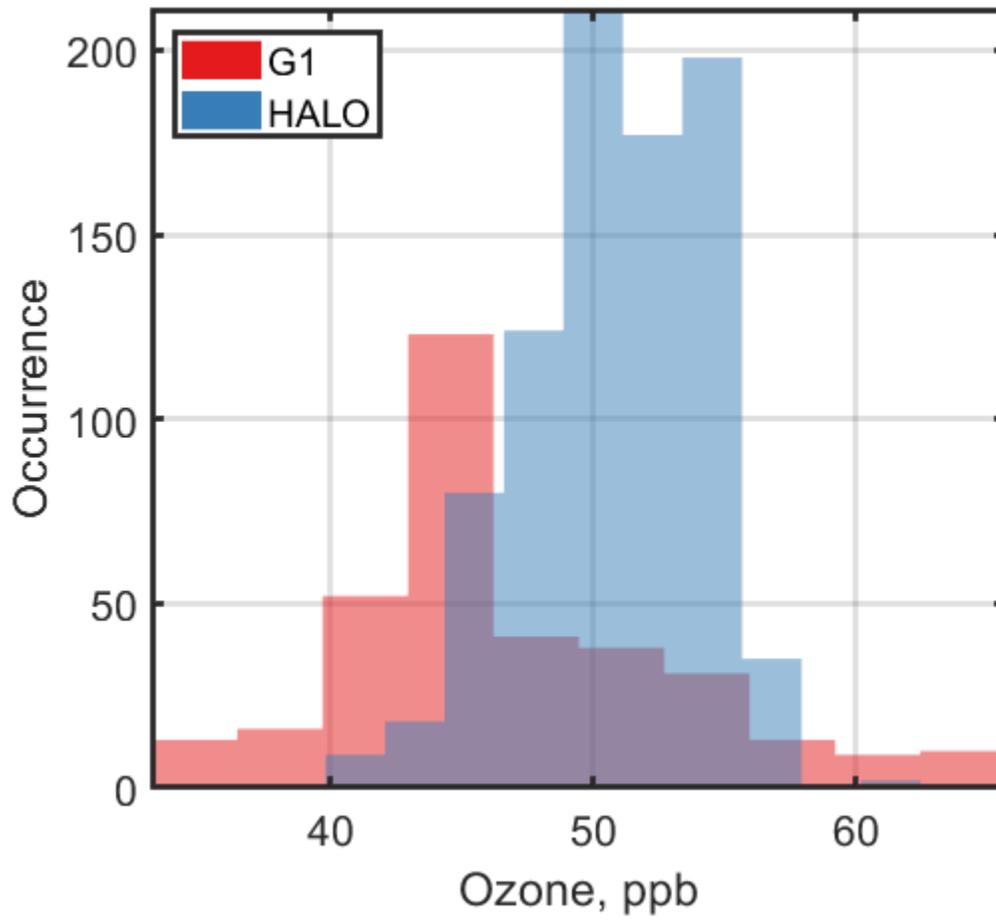
33

34

35 Figure S4. The total aerosol particles number concentration between 2000-3000 m altitude on

36

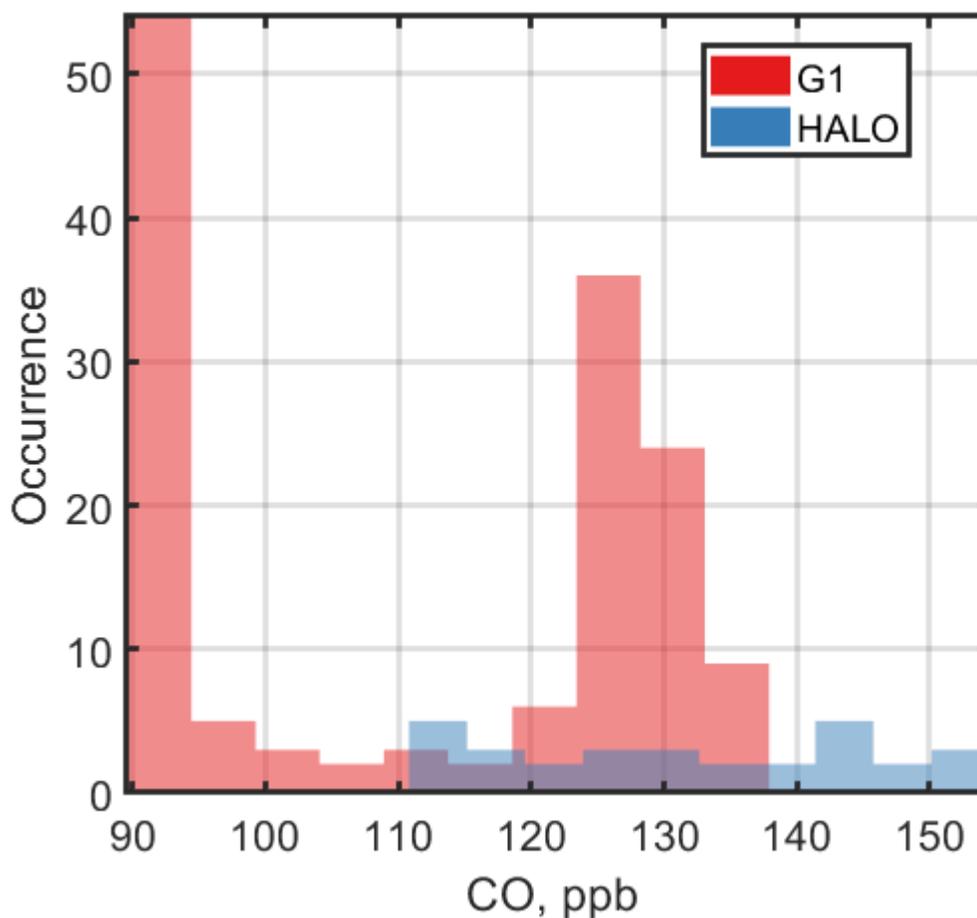
September 21, 2014: (a) CPC measurement; (b) UHSAS measurement.



(a)

37

38



(b)

Figure S5. The trace gas concentration between 2000-3000 m altitude on September 21, 2014:
 (a) Ozone measurement; (b) CO measurement.

2. Additional information for AMS

Most of the details for the AMS measurements have been included in the separate AMS papers (Schulz et al., 2018; Shilling et al., 2018). Brief summaries are provided below.

The G1 AMS was operated with a constant pressure inlet (CPI), which was set to a constant pressure during the campaign. The G1 AMS was calibrated once a week during the deployment. One additional calibration was performed after the flight day, and all the calibrations were in agreement with each other. Based on five calibrations, the averaged parameters such as the

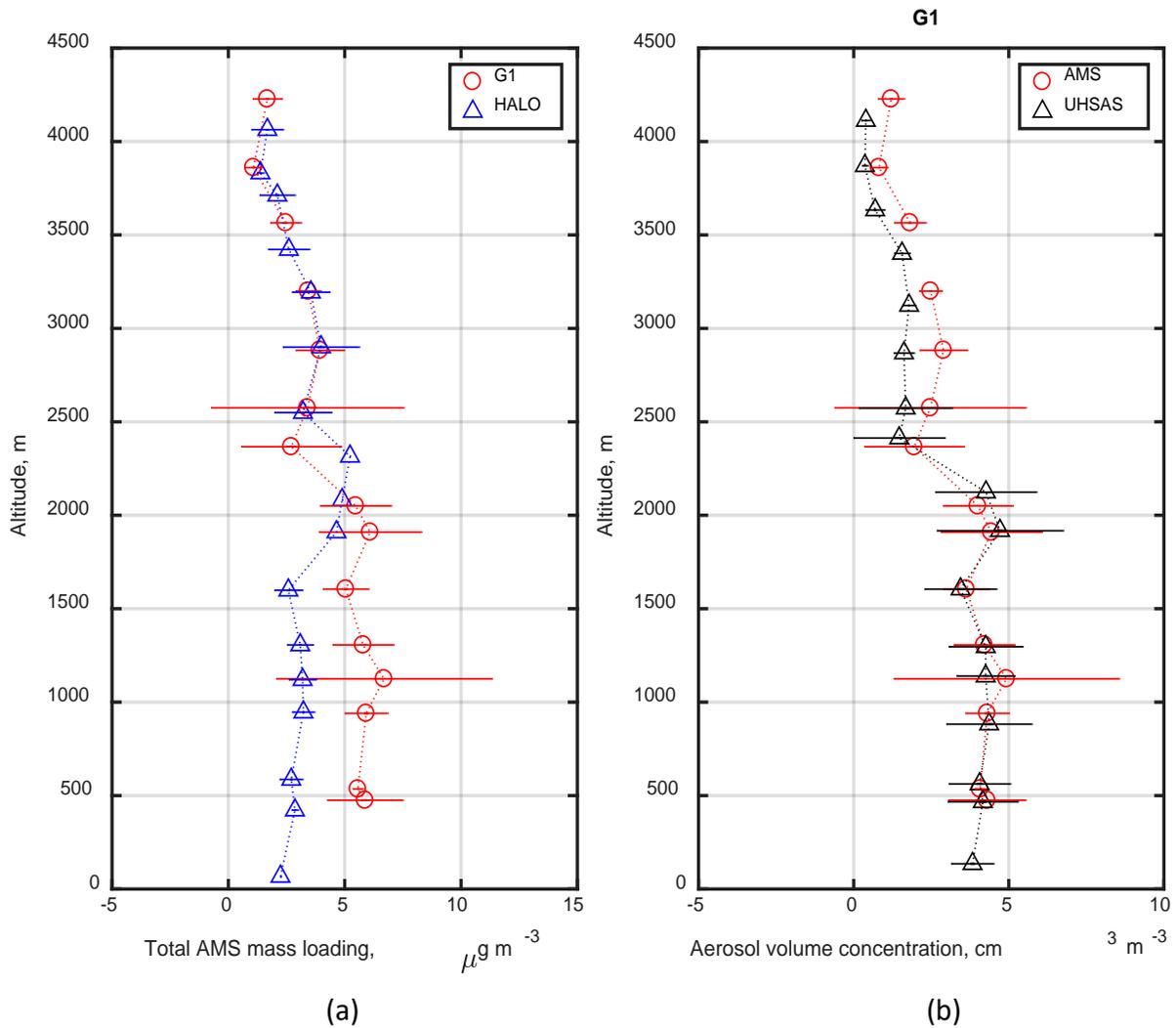
51 airbeam signal (AB), the ionization efficiency (IE), and the relative ionization efficiency (RIE)
52 were applied to all of the data. A real-time correction was made to account for the variations in the
53 AB changes to improve the instrument sensitivity. Typically, this correction is small (<20%) in
54 absolute magnitude. The particle collection efficiency (CE) was determined by comparing AMS
55 data to UHSAS and FIMS data. We also confirmed the CE=0.5 by comparing mass loadings
56 observed at the T3 site to the G1 data.

57 The HALO AMS was calibrated before, during (twice), and after the campaign for (relative)
58 ionization efficiencies of nitrate, ammonium, and sulfate (Schulz et al., 2018). For organics, the
59 default relative ionization efficiency of 1.4 was assumed. The inlet flow was kept constant by the
60 CPI and was measured before and during the campaign. Collection efficiency of 0.5 was applied,
61 as recommended by Middlebrook et al. (2012) for low nitrate conditions. Further details on the
62 operation of the C-ToF-AMS are given in Schulz et al. (2018).

63 Figure S6(a) shows vertical profiles of the total mass concentrations measured by the two
64 AMS instruments on September 21. Above 2500 m altitude, the agreement between the two
65 instruments is excellent (mean difference less than 5%). Between 2000 and 2500 m, the agreement
66 is within the uncertainty range. Below 2000 m altitude, however, the aerosol particle mass
67 concentrations measured by the AMS operated on HALO are lower than the concentrations
68 measured by the AMS on the G1. To compare AMS data to UHSAS data, the aerosol mass
69 concentrations of the G1 AMS were converted to the aerosol volume concentration assuming an
70 organic compound density of 1.5 g cm^{-3} (Pöschl et al., 2010). The converted aerosol volume
71 concentration agreed well with the volume concentration calculated based on UHSAS data,
72 especially below 2500 m, as shown in Figure S6(b). The agreement at lower altitudes suggests that
73 the lower concentration measured by the HALO AMS is due to the transmission efficiency issue
74 in the constant pressure inlet used by the HALO AMS. This inlet was a prototype, designed and
75 built at MPIC Mainz, and works by changing the size of the critical orifice that regulates the flow
76 into the aerodynamic lens. The design and transmission characteristics will be described in an
77 upcoming publication (Molleker, S., in prep.). The AMS aboard the G1 used a constant pressure
78 inlet based on the design in Bahreini et al., 2008. Thus, we conclude that data above 2500 m
79 altitude measured by the AMS aboard HALO in 2014 are valid, while data below 2500 m need to
80 be corrected using correction factors derived from laboratory characterization before further study.

81 After 2014, the HALO AMS inlet design was improved to address the inlet transmission issues
82 specific to this field campaign.

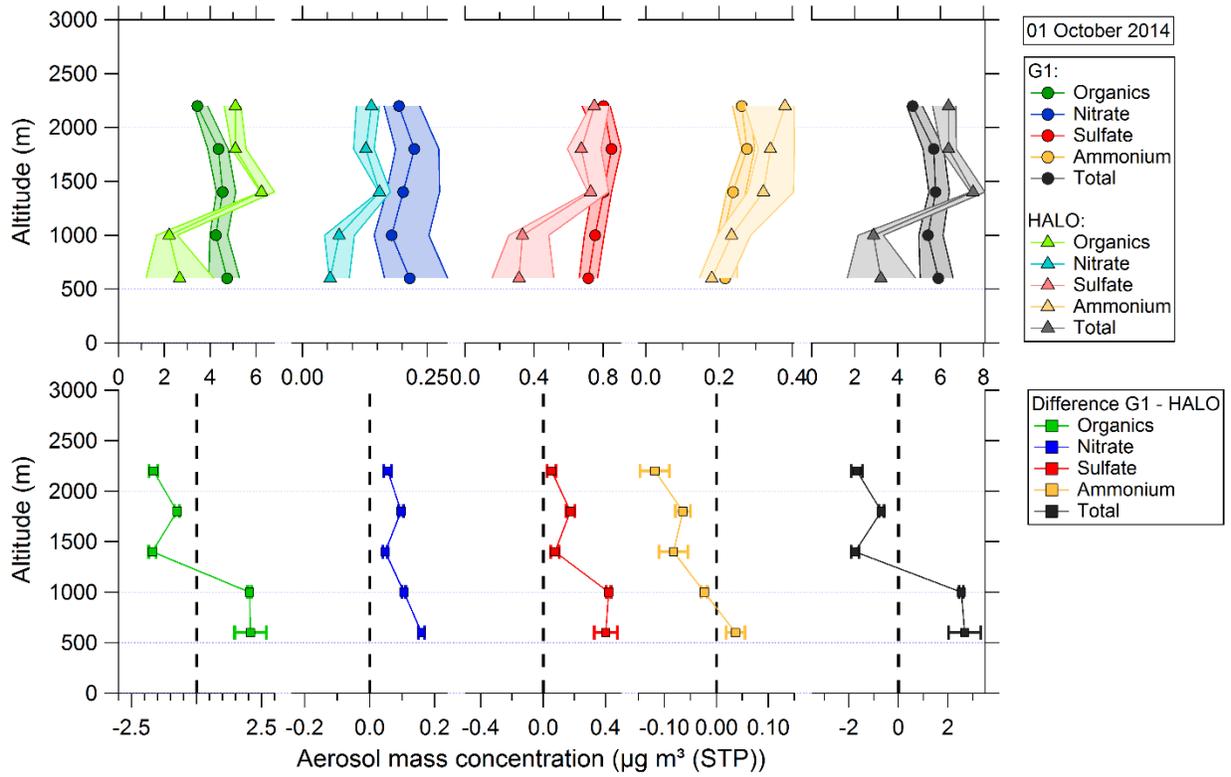
83 The second comparison between the two AMS conducted on October 1 is shown in Figures
84 S6 and S7. The findings are basically in agreement with those of September 21, although the
85 underestimation of aerosol mass concentration due to the inlet in the HALO AMS appears here to
86 be restricted to altitudes lower than 1500 m.



87

88 Figure S6. (a) Comparison of aerosol mass loading measured by the G1 and HALO AMS on
89 September 21; (b) aerosol volume concentration comparison from AMS and the integrated
90 UHSAS on the G1.

91



93

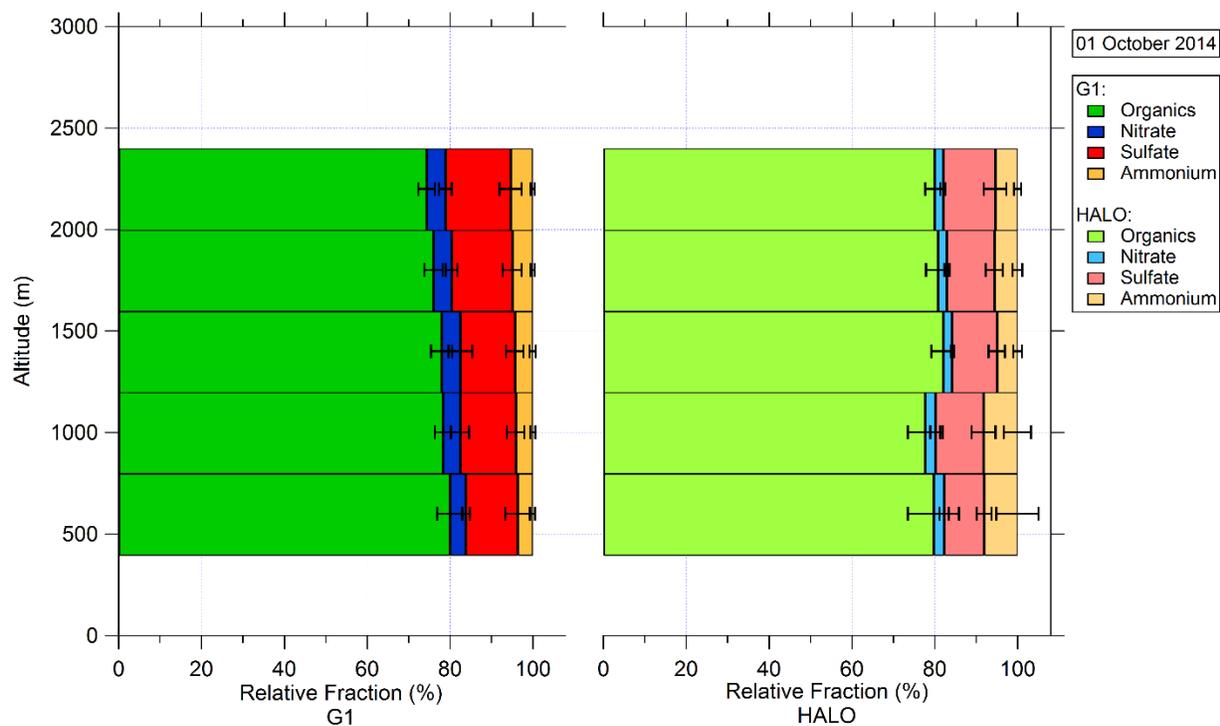
94

Figure S7. The vertical profiling of the aerosol mass concentration observed by the G1 and HALO during October 1.

95

96

97



98
 99 Figure S8. The vertical profiling of the relative fractions for the chemical species observed by the
 100 G1 and HALO during October 1.
 101

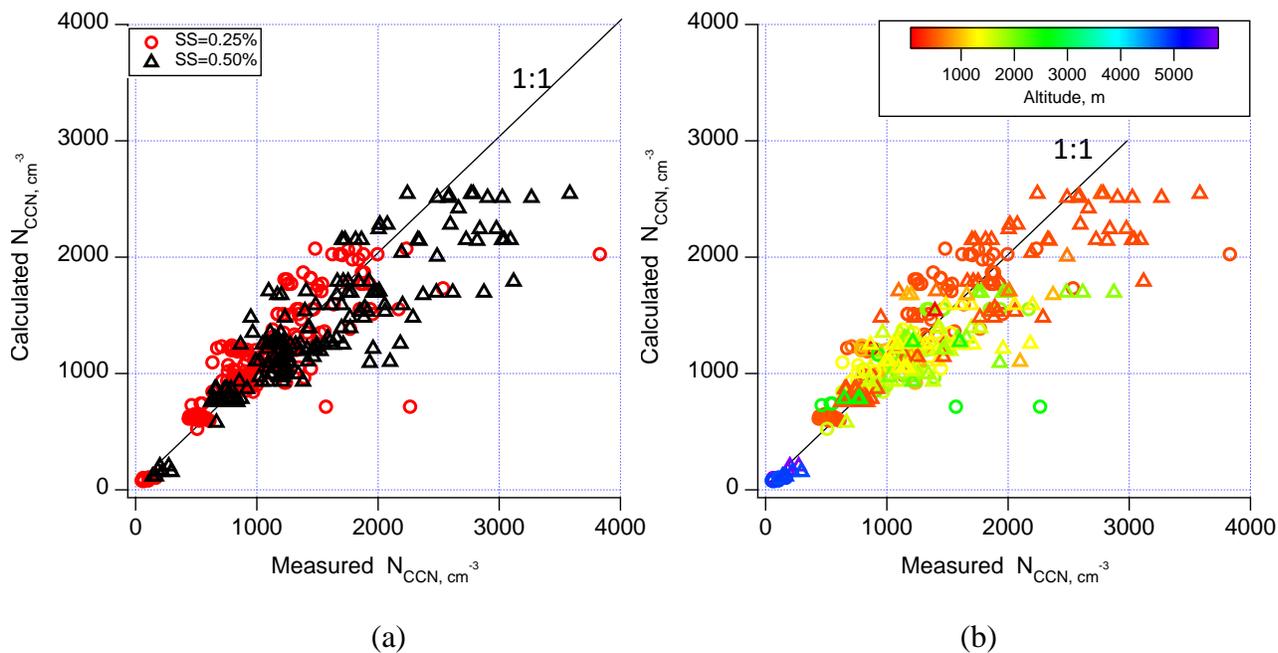
102 3. CCN closure

103 To further examine the relative importance of mixing state and chemical composition, the
 104 CCN concentrations were calculated from aerosol particle size distribution, and chemical
 105 composition measured onboard the G1. The calculation was based on κ -Köhler parameterization,
 106 (Köhler, 1936; Petters and Kreidenweis, 2007, 2008, 2013) and the detail of the approach was
 107 described by Mei et al. (2013b). For the flight on September 9, 2017, the CCN number
 108 concentration calculated from the G1 UHSAS size distribution and chemical composition exhibits
 109 underestimation at a supersaturation of 0.5% (Fig, S9(a)) and when the altitude is below 1000 m
 110 (Fig S9(b)). This underestimation suggests that the UHSAS size range (90-500 nm) did not fully
 111 cover the aerosols with the critical activation diameter ($D_{p,50}$) at high supersaturation. Thus, the
 112 FIMS measurements onboard the G1 was the more appropriate size distribution for both the CCN
 113 closure study. The CCN concentration calculated using the size distribution from FIMS agrees

114 well with the measurement (Fig. S10). The scattering of the comparison data in Figure 15 is likely
115 due to the chemical composition and mixing state effect on aerosol hygroscopicity.

116

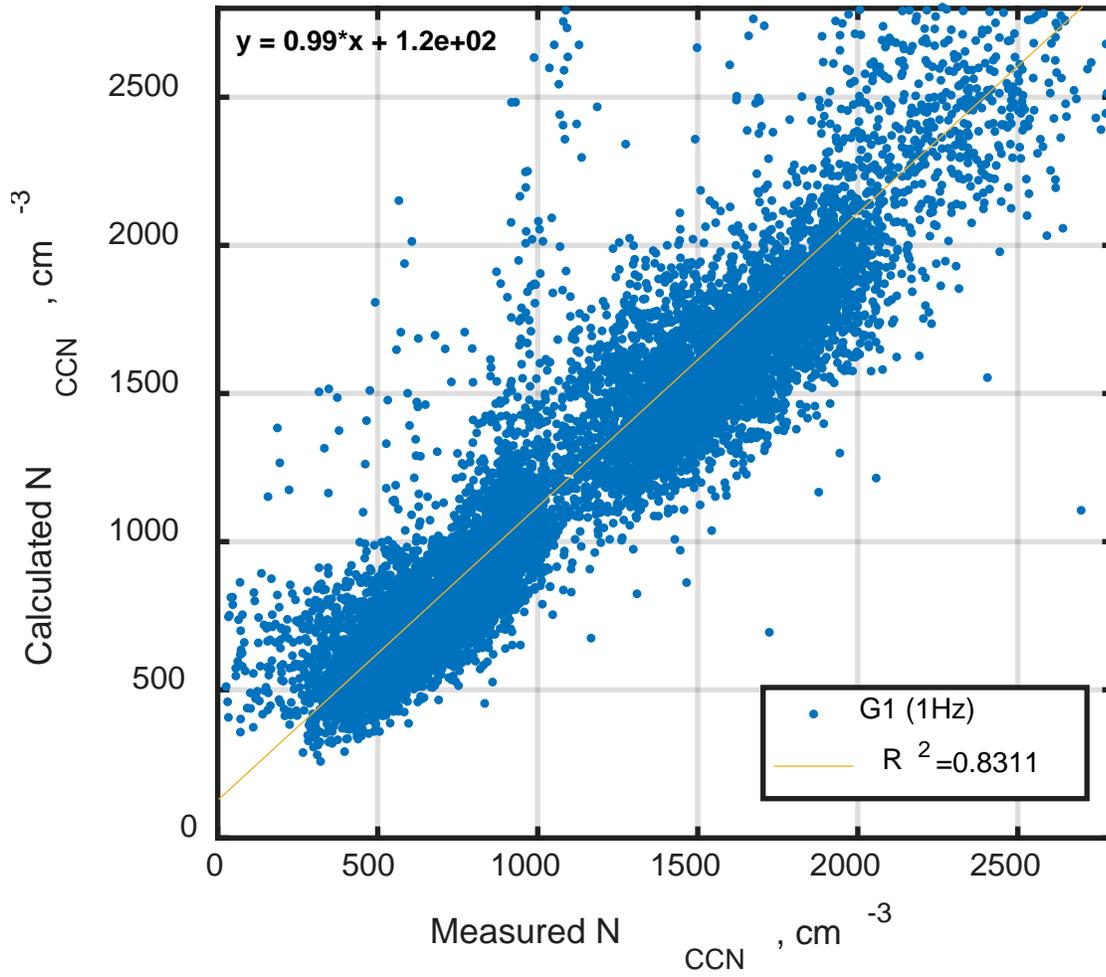
117



118

119

120 Figure S9. Comparison of calculated CCN with measured CCN using the averaged 1 min
121 measurements from the G1: (a) colored by different supersaturations. (b) colored by different
122 altitudes. (Note that both plots used the calculated CCN number concentration from UHSAS size
123 distribution.)



124

125 Figure S10. The scatter plot of the calculated CCN number concentration using FIMS size
 126 distribution compared with the measured CCN number concentration

127

128

129

130

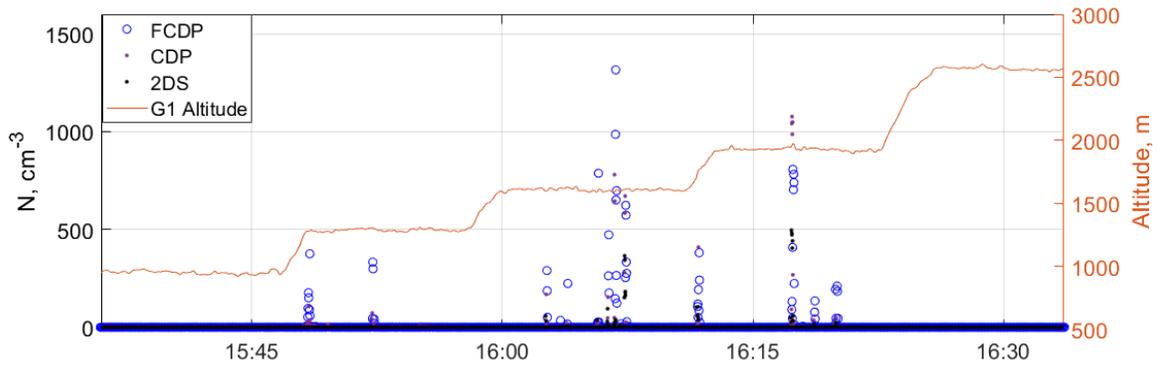
131

132

133

134

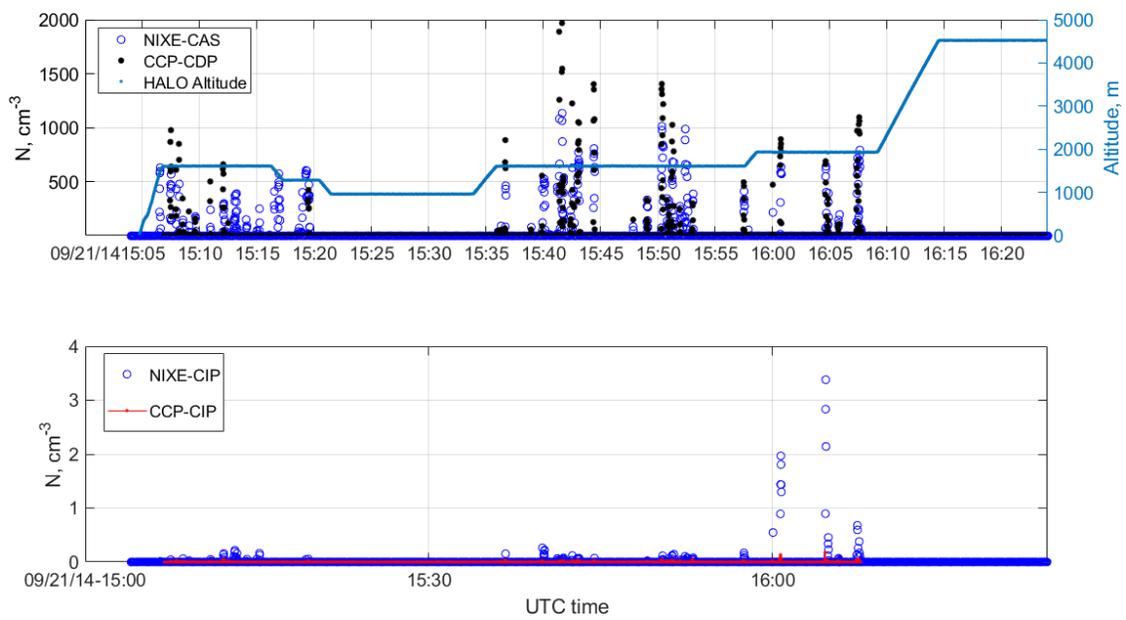
135 **4. Cloud probe observations**



136

137 Figure S11. The cloud droplet number concentration from the G1 aircraft on September 21.

138



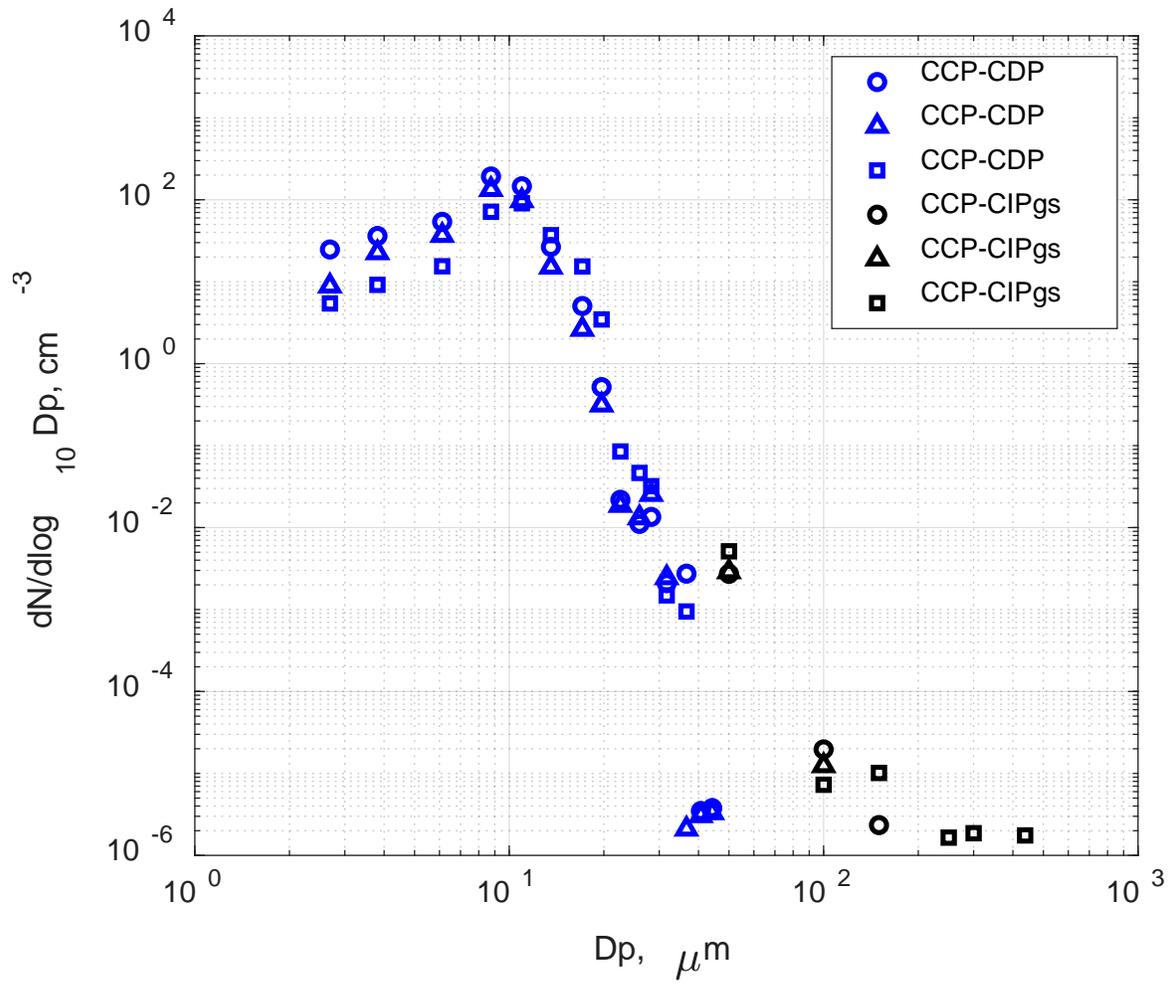
139

140 Figure S12. The cloud droplet number concentration from HALO on September 21.

UTC time: 2014-09-21 15:35:41 - 15:45:19

UTC time: 2014-09-21 15:46:35 - 15:54:46

UTC time: 2014-09-21 15:56:19 - 16:09:33



(a)

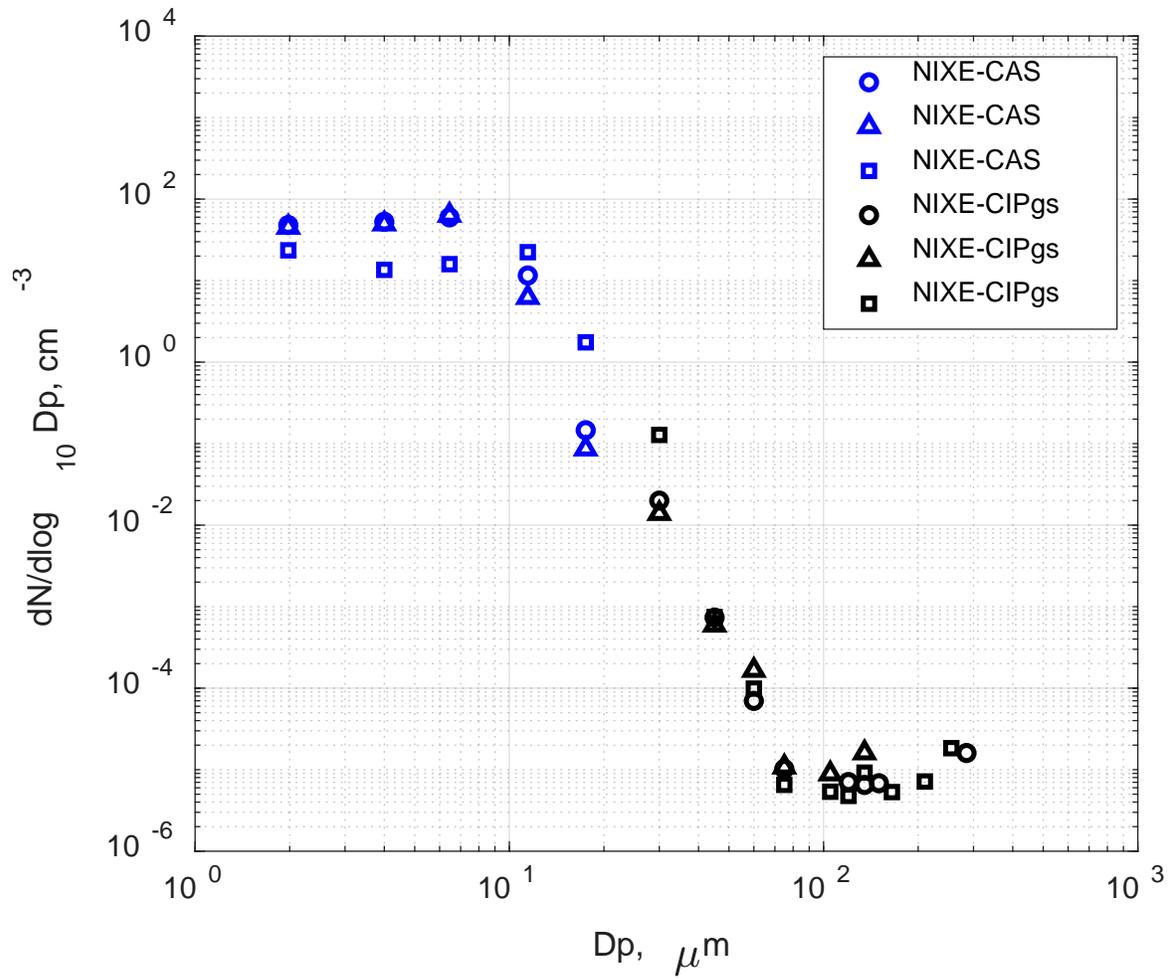
141

142

UTC time: 2014-09-21 15:35:41 - 15:45:19

UTC time: 2014-09-21 15:46:35 - 15:54:46

UTC time: 2014-09-21 15:56:19 - 16:09:33



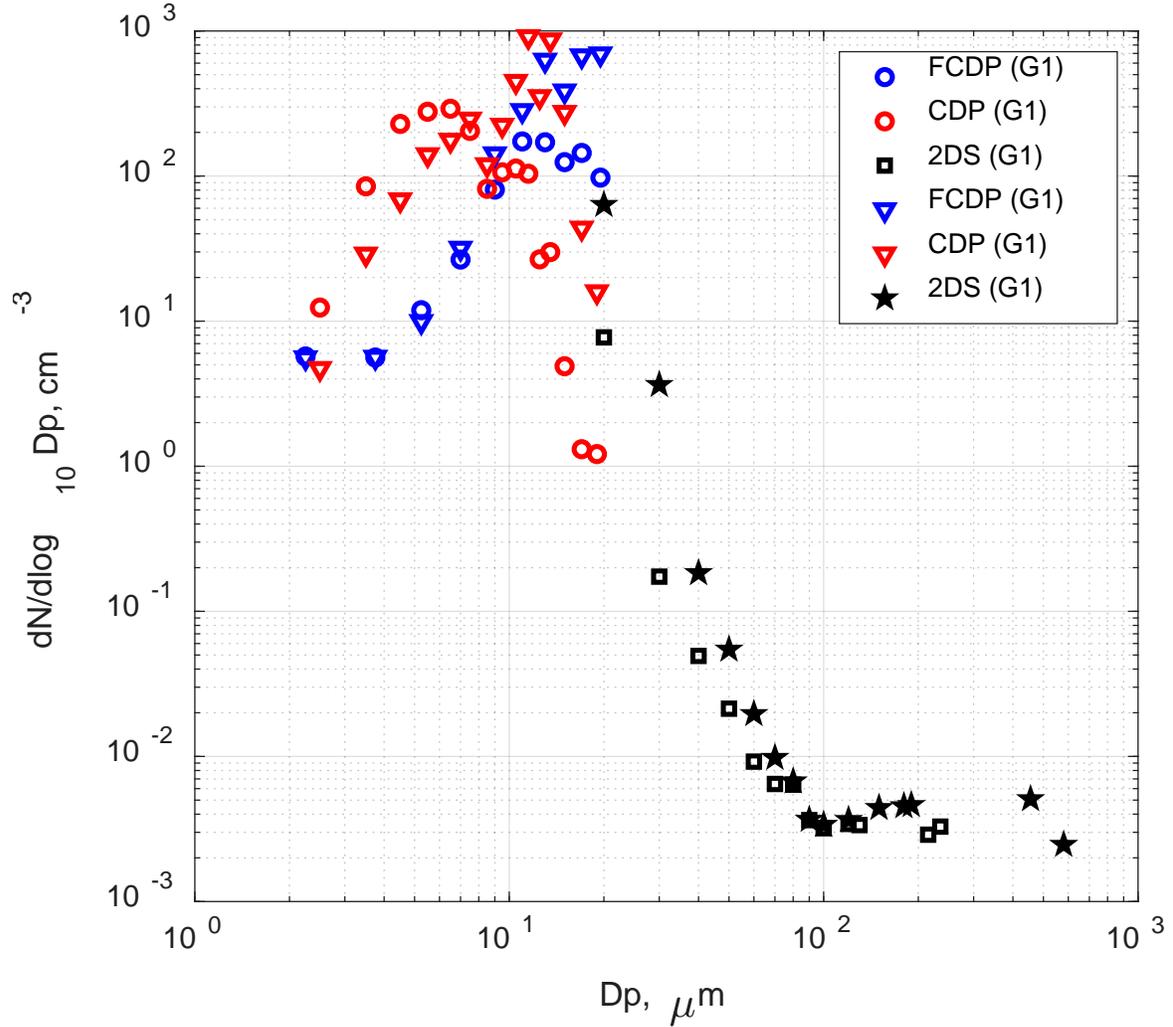
143

144

(b)

UTC time: 2014-09-21 16:02:21 - 16:07:53

UTC time: 2014-09-21 16:16:23 - 16:19:21

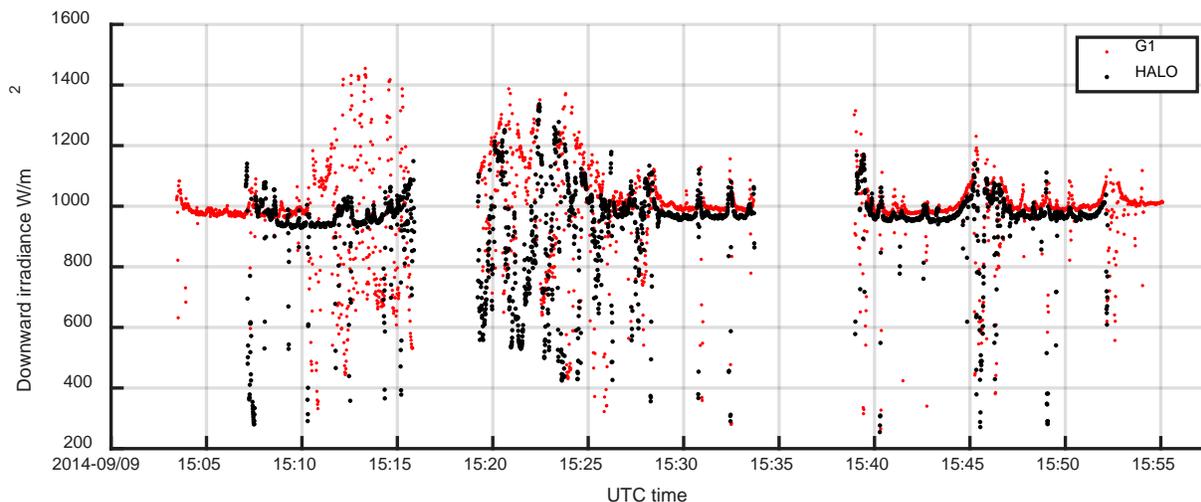


(c)

Figure S13. The averaged cloud droplet size distributions from HALO on September 21, (a) CCP probes; (b) NIXE-CAPS probes; (c) Cloud probes on board the G1.

145
146
147
148
149
150
151
152
153
154

155 **5. Radiation measurements**



156
 157 Figure S14. Time series comparison of the G1 (SPN-1) and HALO (SMART-Albedometer)
 158 radiation measurements on September 9.

159
 160 Table S1. Calibration and maintenance for the instruments deployed on G1

Measurement Variables	Instruments deployed on the G1 (Martin et al., 2016; Schmid et al., 2014)	Calibration/Maintenance
Static Pressure	Rosemount (1201F1), 0-1400 hPa	Calibrated before/after each field campaign
Static air temperature	Rosemount E102AL/510BF -50 to +50 °C	Calibrated before/after each field campaign
Dewpoint temperature	Chilled mirror hygrometer 1011B -40 to +50 °C	Calibrated before/after each field campaign
3-D wind	Aircraft Integrated Meteorological Measurement System 20 (AIMMS-20)	Calibrated with Special flight pattern before each field campaign. Inter-comparison with other GPS/INS during deployment.
Particle number concentration	CPC, cut off size (D_p) =10 nm	Calibrated before/after each field campaign. Weekly calibration of sample and sheath flow rates and inter-comparisons with similar counters during deployment.
Size distribution*	UHSAS-A, 60-1000 nm.	Calibrated before/after each field campaign. Weekly check of sizing with PSL
FIMS	10 nm – 500 nm	Calibrated before/after each field campaign. Weekly calibration of sample and sheath flow rates and checks with one size PSL

Non-Refractory particle chemical composition	HR-ToF-AMS: Organics, Sulfate, Nitrate, Ammonium, Chloride, 60-1000 nm	Weekly calibrations.
CCN concentration	CCN-200, SS= 0.25, 0.5%	Calibrated before/after each field campaign. Biweekly calibration with ammonium sulfate particles.
Gas phase concentration	N ₂ O/CO and Ozone Analyzer, CO, O ₃ concentration, precision 2 ppb	Calibrated before/after each field campaign with calibration gas mixture.
CDP	2-50 μm , $\Delta D_p=1-2 \mu\text{m}$	Calibrated before/after each field campaign by the vendor. Weekly check of sizing with glass beads of several sizes
FCDP	2-50 μm , $\Delta D_p=1-2 \mu\text{m}$	Calibrated before/after each field campaign by the vendor. Weekly check of sizing with glass beads of several sizes
2DS	10-1000 μm	Calibrated before/after each field campaign by the vendor.
Radiation	SPN1 downward irradiance, 400-2700 nm	Calibrated before/after each field campaign

161

162 Table S2. Calibration and maintenance for the instruments deployed on HALO

Measurement Variables	Instruments deployed on HALO (Wendisch et al., 2016)	Calibration/Maintenance
Static Pressure	Instrumented nose boom tray (DLR development), 0-1400 hPa	Calibrated before/after each field campaign
Static air temperature	Total Air Temperature (TAT) inlet (Goodrich/Rosemount type 102) with an open wire resistance temperature sensor (PT100), -70 to +50 °C	Calibrated before/after each field campaign
Dewpoint temperature	Derived from the water-vapor mixing ratio, which is measured by a tunable diode laser (TDL) system (DLR development), 5-40000 ppmv	Calibrated before/after each field campaign
3-D wind	Instrumented nose boom tray (DLR development) with an air data probe (Goodrich/Rosemount) 858AJ and high-precision Inertial Reference System (IGI IMU-IIe)	Calibrated before/after each field campaign
Particle number concentration	CPC, cut off size (D_p) =10 nm	Calibrated before/after each field campaign. Weekly inter-comparisons with similar counters during deployment.
Size distribution*	UHSAS-A, 60-1000 nm.	Calibrated before/after each field campaign. Weekly check of sizing with PSL
Non-Refractory particle chemical composition	C-ToF-AMS: Organics, Sulfate, Nitrate, Ammonium, Chloride, 60-1000 nm	Calibrated before and after the campaign and twice during the campaign

CCN concentration	CCN-200, SS= 0.13-0.53%	Calibrated before/after each field campaign. Weekly calibration with ammonium sulfate particles.
Gas phase concentration	N2O/CO and Ozone Analyzer, CO, O ₃ concentration, precision 2 ppb	Calibrated before/after each field campaign with calibration gas mixture.
Cloud properties*	CCP-CDP, 2.5-46 μm, ΔD _p =1-2 μm	Calibrated before/after each field campaign. Weekly check of sizing with glass beads of several sizes
	NIXE-CAS: 0.61 -52.5 μm	Calibrated before/after each field campaign. Weekly check of sizing with glass beads of several sizes
	NIXE-CIPgs, 15-960 μm	Calibrated before/after each flight with a spinning disk.
	CCP-CIPgs: 15-960 μm	Calibrated before/after each flight with a spinning disk.
Radiation	SMART Albedometer, downward spectral irradiance, 300-2200 nm	Weekly calibrations.

163

164 Table S3. List of compared measurement ranges and measurement variances caused by the spatial
 165 variation during the field campaign.

Measurement Variables	Measured Range during the Field Campaign	Measurement Variances between the Two Aircraft
Static Pressure	500 – 1010 hPa	< 1 %
Static air temperature	272 – 310 K	< 1%
Dewpoint temperature	230 -300 K	Without clouds, <1% With clouds, the measurement from the G1 can be up to 5% lower than that of HALO
3-D wind	1-15 m/s	< 40%
Particle number concentration	500 – 15,000 cm ⁻³	< 20% for CPC, <50% for UHSAS (size dependent)
Non-Refractory particle chemical composition	< 10 μg·m ⁻³	< 10% above 2500 m Up to 50% below 2500 m
CCN concentration	SS=0.25%, 100 – 2000 cm ⁻³	< 10% above 2500 m Up to 50% below 2500 m
Gas phase concentration	Ozone: 15-75 ppb CO: 50-200 ppb	Ozone: < 25% CO: < 15%
Cloud droplet number concentration	3- 20 μm	<50 %
Downward irradiance	200 -1500 W·m ⁻²	< 10%

166

167

168

169 **Reference**

170 Fan, J., Rosenfeld, D., Zhang, Y., Giangrande, S. E., Li, Z., Machado, L. A., Martin, S. T., Yang, Y., Wang, J.,
171 and Artaxo, P.: Substantial convection and precipitation enhancements by ultrafine aerosol particles,
172 *Science*, 359, 411-418, 2018.

173 Kohler, H.: The nucleus in and the growth of hygroscopic droplets, *Transactions of the Faraday Society*,
174 32, 1152-1161, 1936.

175 Kotchenruther, R. A. and Hobbs, P. V.: Humidification factors of aerosols from biomass burning in Brazil, *J*
176 *Geophys Res-Atmos*, 103, 32081-32089, 1998.

177 Moran-Zuloaga, D., Ditas, F., Walters, D., Saturno, J., Brito, J., Carbone, S., Chi, X. G., de Angelis, I. H., Baars,
178 H., Godoi, R. H. M., Heese, B., Holanda, B. A., Lavric, J. V., Martin, S. T., Ming, J., Pohlker, M. L.,
179 Ruckteschler, N., Su, H., Wang, Y. Q., Wang, Q. Q., Wang, Z. B., Weber, B., Wolff, S., Artaxo, P., Poschl, U.,
180 Andreae, M. O., and Pohlker, C.: Long-term study on coarse mode aerosols in the Amazon rain forest with
181 the frequent intrusion of Saharan dust plumes, *Atmos Chem Phys*, 18, 10055-10088, 2018.

182 Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud
183 condensation nucleus activity, *Atmos Chem Phys*, 7, 1961-1971, 2007.

184 Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud
185 condensation nucleus activity - Part 2: Including solubility, *Atmos Chem Phys*, 8, 6273-6279, 2008.

186 Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud
187 condensation nucleus activity - Part 3: Including surfactant partitioning, *Atmos Chem Phys*, 13, 1081-1091,
188 2013.

189 Pöschl, U., Martin, S., Sinha, B., Chen, Q., Gunthe, S., Huffman, J., Borrmann, S., Farmer, D., Garland, R.,
190 and Helas, G.: Rainforest aerosols as biogenic nuclei of clouds and precipitation in the Amazon, *Science*,
191 329, 1513-1516, 2010.

192 Schulz, C., Schneider, J., Amorim Holanda, B., Appel, O., Costa, A., de Sá, S. S., Dreiling, V., Fütterer, D.,
193 Jurkat-Witschas, T., Klimach, T., Krämer, M., Martin, S. T., Mertes, S., Pöhlker, M. L., Sauer, D., Voigt, C.,
194 Weinzierl, B., Ziereis, H., Zöger, M., Andreae, M. O., Artaxo, P., Machado, L. A. T., Pöschl, U., Wendisch,
195 M., and Borrmann, S.: Aircraft-based observations of isoprene epoxydiol-derived secondary organic
196 aerosol (IEPOX-SOA) in the tropical upper troposphere over the Amazon region, *Atmos. Chem. Phys.*
197 *Discuss.*, 2018, 1-32, 2018.

198 Shilling, J. E., Pekour, M. S., Fortner, E. C., Artaxo, P., Sá, S. d., Hubbe, J. M., Longo, K. M., Machado, L. A.,
199 Martin, S. T., and Springston, S. R.: Aircraft observations of the chemical composition and aging of aerosol
200 in the Manaus urban plume during GoAmazon 2014/5, *Atmos Chem Phys*, 18, 10773-10797, 2018.

201 Wang, J., Krejci, R., Giangrande, S., Kuang, C., Barbosa, H. M., Brito, J., Carbone, S., Chi, X., Comstock, J.,
202 Ditas, F., Lavric, J., Manninen, H. E., Mei, F., Moran-Zuloaga, D., Pohlker, C., Pohlker, M. L., Saturno, J.,
203 Schmid, B., Souza, R. A., Springston, S. R., Tomlinson, J. M., Toto, T., Walter, D., Wimmer, D., Smith, J. N.,
204 Kulmala, M., Machado, L. A., Artaxo, P., Andreae, M. O., Petaja, T., and Martin, S. T.: Amazon boundary
205 layer aerosol concentration sustained by vertical transport during rainfall, *Nature*, 539, 416-419, 2016.

206 Williamson, C., Kupc, A., Wilson, J., Gesler, D. W., Reeves, J. M., Erdesz, F., McLaughlin, R., and Brock, C.
207 A.: Fast time response measurements of particle size distributions in the 3-60 nm size range with the
208 nucleation mode aerosol size spectrometer, *Atmos Meas Tech*, 11, 3491-3509, 2018.